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Integrated geothermal-CO₂ reservoir systems: Reducing carbon intensity through sustainable energy production and secure CO₂ storage

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Abstract

Large-scale geologic CO₂ storage (GCS) can be limited by overpressure, while geothermal energy production is often limited by pressure depletion. We investigate how synergistic integration of these complementary systems may enhance the viability of GCS by relieving overpressure, which reduces pore-space competition, the Area of Review, and the risks of CO₂ leakage and induced seismicity, and by producing geothermal energy and water, which can defray parasitic energy and water costs of CO₂ capture.

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1. Introduction

Geothermal energy and geologic CO₂ storage (GCS) have the potential of significantly contributing to lowering atmospheric CO₂ emissions, which is necessary for mitigating climate change [1,2]. For large-scale CO₂ injection in saline formations, overpressure can limit the ability to store CO₂, while geothermal energy production can be limited by pressure depletion [3,4]. These two complementary systems can be integrated synergistically, with CO₂ injection providing pressure support to maintain the productivity of geothermal wells, while the net loss of brine provides pressure relief and improved injectivity for CO₂-injection wells. The loss of brine can occur as an outcome of geothermal operations [4] and/or be engineered through the consumptive beneficial use of brine for either freshwater generation or for saline cooling purposes [5,6]. Relief of overpressure can increase CO₂ storage capacity, reduce the risks of induced seismicity, as well as decrease CO₂ and brine leakage [6,7]. Depending on reservoir conditions, injected CO₂ can also be utilized as a more efficient working fluid (than brine) for geothermal heat extraction [8-10]. We examine a range of pressure-management strategies for integrating GCS with the production of geothermal energy and water in permeable saline aquifers.

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2. Background

For large-scale deployment of GCS in saline aquifers to be achievable, several implementation barriers must be addressed. CO₂ capture costs, sequestration safety, and public acceptance have been recognized for a number of years [1]. Receiving more recent attention are water-use demands from CO₂ capture, pore-space competition with emerging activities, such as shale-gas production, and induced seismicity [11,12]. Besides parasitic energy and water costs associated with CO₂ capture, the primary technical driver for the most challenging implementation barriers is overpressure caused by CO₂ injection.

Utilization of CO₂ to generate economic benefits is motivated by the need to defray CO₂ capture and sequestration costs. Compared to the several decades of experience with CO₂-enabled enhanced oil recovery (EOR), utilization of CO₂ in geothermal energy production has only recently been investigated [3,4,8-10]. Reservoir studies of integrated geothermal-GCS fall under two end-member approaches to CO₂ utilization. The first approach has emphasized utilization of CO₂ as an efficient working fluid because of its advantageous thermo-physical properties, namely the low viscosity and thermosyphon effect of CO₂, which reduce the parasitic power consumption of the working-fluid recirculation system [8-10]. Those studies have focused on small injector/producer well spacing in order to promote early CO₂ breakthrough and recirculation. Because the goal of the first approach is to maximize the heat-extraction benefit per ton of delivered (captured) CO₂, GCS is an ancillary, rather than a primary benefit. The second approach has emphasized utilizing CO₂ injection as a means of providing pressure support for geothermal production wells, rather than utilizing CO₂ as a working fluid [3,4]. For the second approach it is important to delay CO₂ breakthrough at the brine producers in order to maximize their useful lifetime. Hence, those studies have emphasized large injector/producer spacing. A key objective is for brine production, together with its consumptive beneficial use, to provide pressure relief for CO₂ injection, thereby increasing CO₂ storage capacity and efficiency, while reducing (1) pore-space competition with neighboring subsurface activities, (2) the Area of Review (AoR), and (3) environmental risks associated with overpressure [6,7]. A significant ancillary benefit of the second approach is that it can generate large quantities of product water (freshwater of saline cooling water) and thereby defray the parasitic water demand of CO₂ capture.

Ideal settings for GCS are deep sedimentary formations, such as sandstones, which are highly porous and permeable and laterally extensive, overlain by an impermeable laterally-extensive caprock [12]. If temperatures are high enough, such settings are also ideal for integrated geothermal-GCS. Such settings are also well suited for geothermal energy production because the caprock prevents cooler overlying water from being drawn down to the producers [3,4] and laterally extensive, highly permeable formations are conducive to high well productivity and lower hydraulic drawdown, which allows for greater spacing between producers and injectors.

3. Model description

In this study we used the NUFT code, which simulates multi-phase heat and mass flow and reactive transport in porous media [13]. A 3-D model with quarter symmetry is used to represent a 250-m-thick storage aquifer (Fig. 1), with a permeability of 10^{-13} m², similar to previous studies [4]. The bottom of the storage aquifer is located either 2.5 or 5.0 km below the ground surface, and is bounded by impermeable (10^{-18} m²) (caprock and bedrock) *seal* units. The bedrock, which extends to the lower boundary 1 km below the bottom of the storage aquifer, has a no-mass-flow condition and a specific geothermal heat flux of either 75 or 100 mW/m². Above the storage aquifer is a sequence that includes the lower caprock (250 m thick), the middle (saline) aquifer (250 m thick), the upper caprock (250 m thick), and the upper (saline) aquifer, which extends to the ground surface (Fig. 1). Note that we did not discretely represent shallow freshwater aquifers. We also considered cases where the middle aquifer is overlain by a caprock that extends to the ground surface. Overpressure in the storage aquifer and lower caprock was found to be insensitive to whether the upper caprock is 250 m thick or extends all the way to the ground surface. A thermal conductivity of 2.0 W/m °C is applied throughout the model, resulting in thermal

gradients of 37.5 and 50 °C/km. The outer boundaries have a no-flow condition to represent a semi-closed reservoir with an area of 31,416 km². Additional model details are given elsewhere [4].

We modeled symmetrical patterns of horizontal wells, oriented orthogonal to the cross sections in Figs. 1 and 2. For all cases (Table 1), there are two horizontal CO₂ injectors at the bottom of the storage aquifer, each with a perforated length of 4 km. The total CO₂ injection rate is 480 kg/sec (15.15 million tonnes/yr) for 100 years, which is the CO₂ generated by 1.9 GWe of coal power plants. For the large well-spacing cases, CO₂ injectors are spaced 7 km from the center of injection. We also considered small well-spacing cases with CO₂ injectors spaced 2 km from the center of injection. In addition to the horizontal-well injection-only case (I-H), we considered a corresponding case (I-V) with two vertical CO₂ injectors. For the large well-spacing cases with brine production (Fig. 1b and d), four producers are spaced 2 km from the center of injection (5 km from the CO₂ injectors) at the bottom of the storage aquifer. For the small well-spacing cases, the producers are 1 km from the center of injection (1 km from the CO₂ injectors). Because fluid production is limited by the capacity of submersible pumps, we used four producers, each with a production rate of 120 kg/sec and a perforated length of 1 km. Initially, production is limited to brine; however, as CO₂ breaks through, CO₂ and brine are co-produced, with the CO₂ cut (fraction) increasing with time (Fig. 3a). Total fluid production rate (brine plus CO₂) is fixed at 120 kg/sec per well. In addition to pressure management reasons, this well pattern is motivated by the need to centralize fluid production, which minimizes the distance that hot fluids are conveyed to the geothermal plant. All produced fluid is reinjected, with produced brine injected in the brine injectors (Fig. 1b and d), and produced CO₂ returned to the CO₂ injectors. CO₂ and produced brine are injected at a fluid enthalpies corresponding to 16 °C at injection conditions, approximating average annual surface temperatures.

Table 1. Large (5-km) CO₂-injector-producer spacing cases are summarized. Small (1-km) CO₂-injector-producer spacing was also applied to B-3.

Case	Production/ injection mass ratio	Net mass injection rate (kg/sec)	Product water generation rate (acre-ft/yr)	CO ₂ -injector-to- brine-injector spacing	Brine injectors in overlying saline aquifer	Brine-injection- to-production mass ratio	Number of brine producers
I-V	0	480	N/A	N/A	N/A	N/A	N/A
I-H	0	480	N/A	N/A	N/A	N/A	N/A
B-1	1	480	0.0	3 km	0	1.0	0
B-2	1	480	0.0	3 km	2	1.0	0
B-3	1	480	0.0	6 km	2	1.0	0
B-4	1	316.8	4,175	3 km	2	0.66	0
3B-1	3	-9.6	12,526	3 km	4	0.66	8
3B-2	3	-9.6	12,526	3 km	4	0.66	8
4B-1	4	-172.8	16,701	3 km	4	0.66	12

Case B-1 injects all produced brine into two brine injectors at the top of the storage aquifer (Fig. 1b and d). Cases B-2, B-3, B-4, 3B-1, 3B-2, and 4B-1 also include two brine injectors at the bottom of the overlying (middle) aquifer (Fig. 2a and c), where half the produced brine is injected. For cases B-4, 3B-1, 3B-2, and 4B-1, 34 percent of the produced brine is consumed for beneficial use, such as reverse osmosis desalination or saline cooling water, and the remaining 66 percent (residual brine) injected into the storage and middle aquifers. For cases 3B-1, 3B-2, and 4B-1, a total of eight *perimeter* brine producers (each at a rate of 120 kg/sec) are added at the bottom of the middle and storage aquifers, 5 km from the brine injectors (Fig. 2b and d). The purpose of the additional brine producers is to reduce net fluid injection, which reduces overpressure in the storage aquifer and lower caprock, as well as in the far-field. For cases 3B-1, 3B-2, and 4B-1, two brine injectors are added at the bottom of the middle aquifer, above the central producers, in order to more evenly distribute brine injection. Case 3B-2 is modified from 3B-1 by injecting two-thirds (rather than half) of the residual brine into the middle aquifer and the remaining one-third into the storage aquifer. Case 4B-1 is modified from 3B-1 by adding four perimeter brine producers (each at a rate of 120 kg/sec) at the bottom of the storage aquifer (in line with the other brine producers) and by injecting 60 percent (rather than half) of the residual brine into the middle aquifer and 40 percent into the storage aquifer.

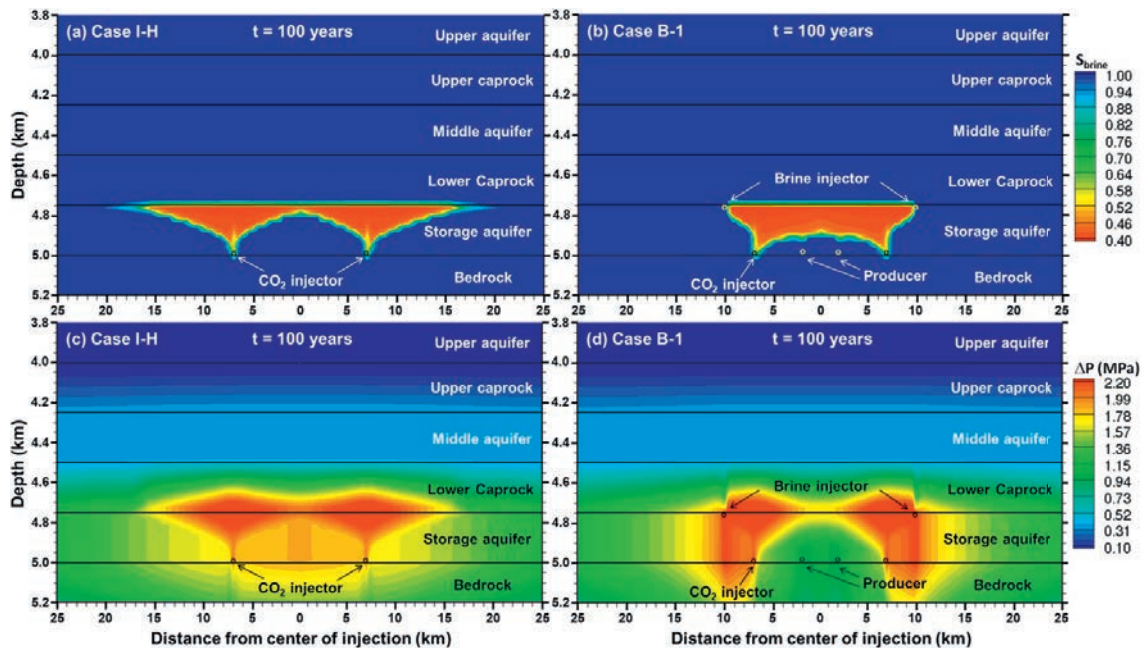


Fig.1 Brine saturation (a,b) and overpressure (c,d) for cases I-H and B-1 and a specific geothermal heat flux of 100 mw/m^2 .

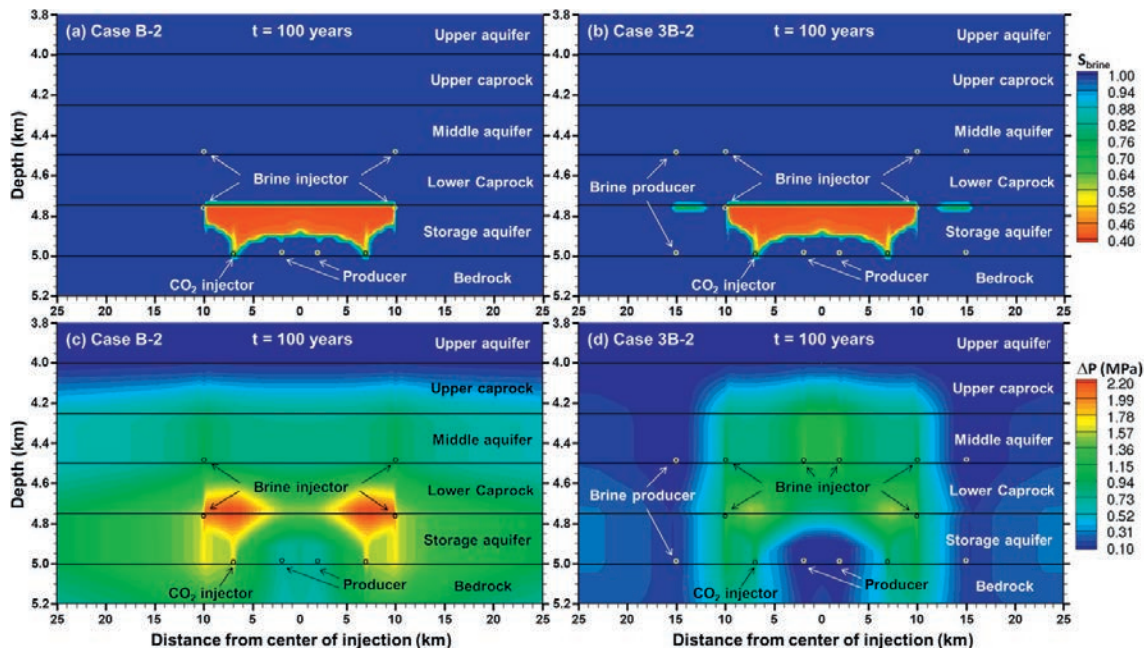


Fig. 2 Brine saturation (a,b) and overpressure (c,d) for cases B-2 and 3B-2 and a specific geothermal heat flux of 100 mw/m^2 .

4. Results

Brine production and injection can achieve several useful objectives: (1) control of the CO₂ plume, (2) vertically extending and shaping the region of overpressure, which reduces the magnitude of overpressure (Fig. 4) and creates a hydraulic barrier to CO₂ leakage, (3) reducing the overpressure difference across the lower caprock (Fig. 5),

which also reduces the potential for CO₂ leakage, (4) reducing far-field overpressure (Fig. 6), and (5) producing large quantities of brine for generating water and electricity.

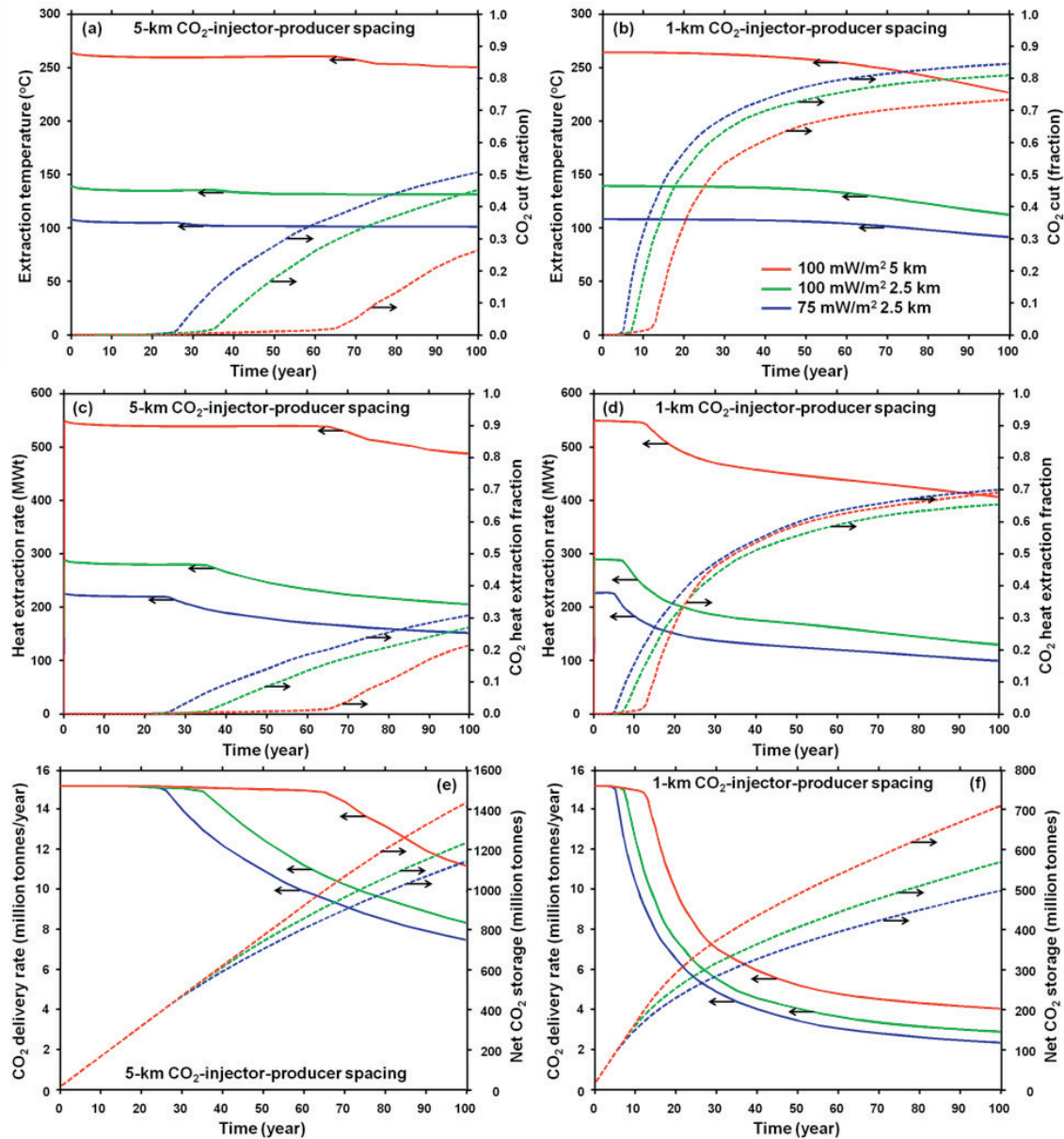


Fig. 3. Histories of extraction temperature and CO₂ cut (a,b); heat extraction rate and CO₂ heat extraction fraction (c,d); and CO₂ delivery rate and net CO₂ storage (e,f) are plotted for case B-2 for large (a,c,e) and small (b,d,f) CO₂-injector-producer spacing.

4.1. Manipulation and control of the CO₂ plume and overpressure

Migration of injected CO₂ is dominated by buoyancy (Fig. 1a). Buoyant flow of CO₂ also dominates where maximum overpressure occurs (Fig. 1c). Rather than occurring at the point of injection, the maximum overpressure occurs above the CO₂ injectors, at the contact between the storage aquifer and the lower caprock. The addition of two

producers pulls the CO₂ plumes towards the center of injection, where they coalesce (Fig. 1b), while the brine injectors create hydraulic barriers (Fig. 1d) that push the CO₂ plumes towards the center of injection. Pull-push manipulation causes the CO₂ plumes to occupy more of the vertical extent of the storage aquifer, and constrains outward migration, with greater utilization of the aquifer pore space. Overpressure in the middle aquifer is identical for these two cases (Fig. 1c and d). The histories of overpressure at the top of the storage aquifer (Fig. 4c), overpressure difference across the caprock (Fig. 5c), and far-field overpressure (Fig. 6c) are the same for these two cases (I-V and B-1). The primary purpose of producing brine and injecting it back into the storage aquifer is to control CO₂ plume migration; however, overpressure is unchanged. A comparison of case I-H with the vertical-well case (I-V) shows the benefit of using horizontal wells to reduce overpressure (Figs. 4, 5, and 6).

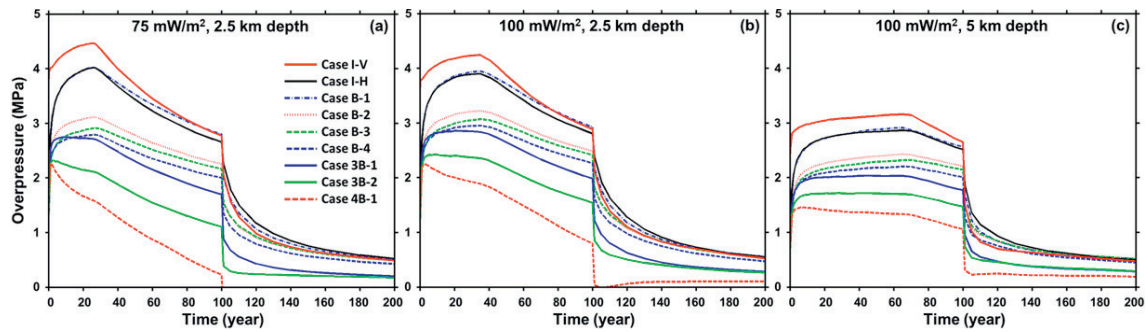


Fig. 4. Histories of maximum overpressure at the top of the storage aquifer above the CO₂ injector are plotted.

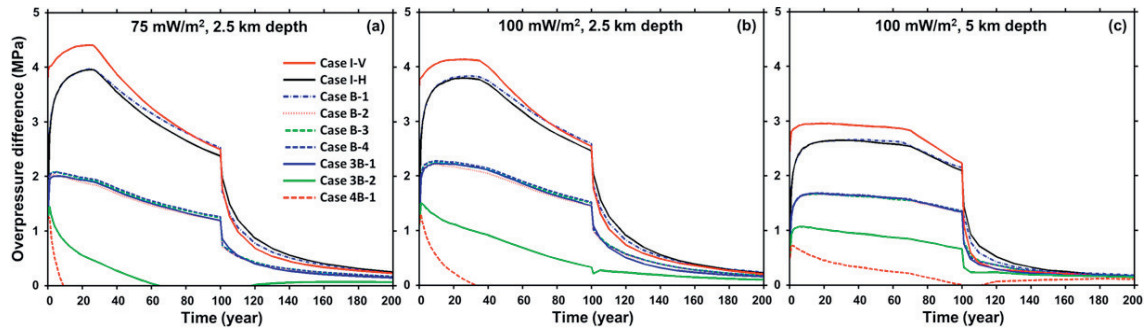


Fig. 5. Histories of overpressure difference across the caprock above the CO₂ injector are plotted.

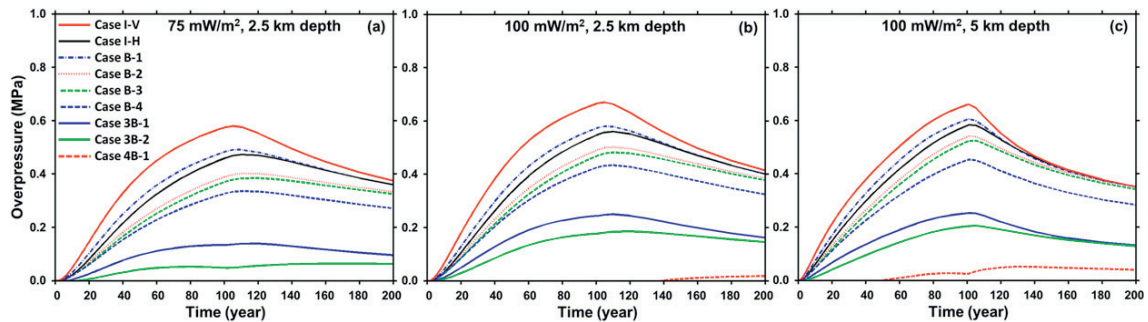


Fig. 6. Histories of overpressure at the top of the storage aquifer, 79 km from the center of injection are plotted.

A comparison of panels a, b, and c in Figs. 4 and 5 illustrates the temperature dependence of overpressure, driven by two different effects. First, hydraulic conductivity increases with temperature, due to the corresponding decrease in viscosity, causing overpressure to decrease with increasing temperature. Second, the viscosity of water decreases more steeply with increasing temperature than does that of supercritical CO₂, causing the mobility ratio

of CO₂ and brine to decrease with temperature. Consequently, preferential CO₂ flow is less pronounced at higher temperatures, which increases CO₂ breakthrough time, causing CO₂ cut to increase more slowly (Fig. 3a). Since the produced CO₂ is recycled, CO₂ delivery rate decreases more quickly for lower temperatures, resulting in less net CO₂ storage (Fig. 3e) and a more rapid decline in overpressure (Figs. 4 and 5). Overpressure peaks when CO₂ first arrives at the producers; thereafter, it declines due to the decreasing CO₂ delivery rate. Because the injection-only cases apply the CO₂ delivery rate history of the corresponding brine production cases, overpressure peaks at the same time as in the brine production cases. Because net CO₂ storage increases with temperature, far-field overpressure increases with temperature (Fig. 6).

Injecting half of the produced brine into the middle aquifer extends the region of overpressure all the way up to the upper caprock, reducing its magnitude in the storage aquifer and lower caprock (Fig. 2); yet, the hydraulic barrier still constrains outward CO₂ migration (Fig. 2). Distributing brine injection over the middle and storage aquifers reduces overpressure at the top of the storage aquifer (Fig. 4). Increasing overpressure in the middle aquifer substantially reduces the overpressure difference across the caprock (Fig. 5c), which reduces the driving force for potential CO₂ leakage through discrete pathways such as abandoned wells or faults. Brine injection in the middle aquifer has a minor effect on far-field overpressure (Fig. 6c) because far-field overpressure is sensitive to net fluid injection. Increasing brine-injector-to-CO₂-injector spacing from 3 to 6 km has a minor impact on decreasing overpressure (cases B-2 and B-3 in Fig. 4). Decreasing the brine injection rate by 34 percent has a minor impact on decreasing overpressure (cases B-3 and B-4 in Fig. 4). Overpressure difference (Fig. 5) is more sensitive to the relative rate of brine injection in the middle and storage aquifers than to the magnitude of those rates.

The most influential brine injection/production strategy is to combine the consumptive use of brine with production/injection ratios > 1, as done in cases 3B-1, 3B-2, and 4B-1, which requires the addition of brine producers, located at the perimeter and spaced 5 km from the brine injectors (Fig. 2b and d). The additional brine producers have little influence on the shape of the CO₂ plume (Figs. 2a and b). The brine injectors create a hydraulic *ridge*, which is a barrier to CO₂ migration and leakage, while the brine producers create a hydraulic *trough*. With high enough production rates, this trough can nullify the hydraulic ridge, causing far-field overpressures to be close to zero (Figs. 2d and 6). This strategy can hydraulically isolate a geothermal-GCS operation from neighboring subsurface activities, which equates to reducing the AoR and pore-space competition. This strategy can also greatly reduce the overpressure difference across the caprock (Fig. 5).

4.2. Production of geothermal energy and water

For all cases, an initial decline in temperature of 3 to 4 °C is the result of cooler water being drawn down to the producer at the base of the storage aquifer (Fig. 3a and b). After this initial period of thermal mixing, thermal drawdown does not commence until the arrival of CO₂ (Fig. 3a and b). We considered cases representative of the two end-member strategies for integrated geothermal-GCS, starting with the strategy that utilizes CO₂ injection for pressure support of geothermal wells. For large (5-km) CO₂-injector-producer spacing, thermal drawdown is small (4, 4, and 10 °C) during the 100-year operation period. Heat extraction rate is highest (221, 282, and 541 MWt) when only brine is produced (Fig. 3c), then declines with increasing CO₂ cut because CO₂ carries less heat per unit mass than brine. Small (1-km) CO₂-injector-producer spacing causes early CO₂ breakthrough (Fig. 3b) and recirculation, with a more rapid decline in CO₂ delivery rate (Fig. 3f). Because CO₂ breakthrough occurs earlier, thermal drawdown is faster (Fig. 3b). Heat extraction rate also declines more steeply. Had we taken advantage of the lower viscosity and thermosyphon effect of CO₂, we could have increased the CO₂ production rate to maintain a higher heat extraction rate than shown in Fig. 3d. For cases 3B-1, 3B-2, and 4B-1, thermal breakthrough does not occur at the brine producers during the 100-year operation period, resulting in negligible thermal drawdown. For case 4B-1, total brine production is 1920 kg/sec, resulting in a heat extraction rate of 842, 1072, and 2115 MWt for the low, medium, and high temperature cases. Using GETEM [14] to calculate binary-cycle electrical power production, case 4B-1 is found to generate 35, 62, and 253 MWe for these cases. For the high-temperature case, this power output is 13.3 percent of the 1.9 GWe output from the coal plants that generated the injected CO₂. Annual water production for these cases is listed in Table 1.

5. Conclusions

We have shown how the viability of large-scale GCS may be enhanced through its synergistic integration with geothermal energy and water production. The most challenging implementation barriers facing GCS are energy and water demands for CO₂ capture and overpressure caused by CO₂ injection. We conduct reservoir analyses of horizontal wells in a stack of permeable saline aquifers, separated by impermeable seal units. We analyze the two end-member approaches to utilizing CO₂ in geothermal energy production: one focused on utilizing CO₂ as an efficient working fluid, and the other focused on utilizing CO₂ injection for pressure support of geothermal wells. For the second approach, we consider different pressure-management strategies and develop a generic well pattern with four concentric zones of wells in the storage aquifer: (1) inner brine/CO₂ producers, (2) CO₂ injectors, (3) brine injectors, and (4) outer brine producers; and two zones (3 and 4) in the overlying saline aquifer. This well pattern manipulates the CO₂ plume by pulling and pushing, creates a hydraulic ridge and cap to suppress CO₂ migration and leakage, and creates an outer hydraulic trough to nullify the hydraulic ridge, thereby hydraulically isolating the geothermal-GCS operation from its neighbours, which reduces pore-space competition and the Area of Review (AoR). In addition to managing overpressure, we show how this concept has the potential of producing enough geothermal energy and water to significantly defray the parasitic costs of CO₂ capture.

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